

GEWEX Process Evaluation Study on Upper Tropospheric Clouds & Convection

GEWEX UTCC PROES

advance understanding on feedback of UT clouds

large-scale modelling necessary to identify most influential feedback mechanisms

-> models should be in agreement with observations

- *understand relation between convection, cirrus anvils & radiative heating*
- *provide observational metrics to probe process understanding*

Claudia Stubenrauch

Laboratoire de Météorologie Dynamique / IPSL, France



GEWEX Data and Assessment Panel meeting

9 – 12 Oct 2017, Boulder, USA

GEWEX UTCC PROES participants

Coordination: Claudia Stubenrauch & Graeme Stephens

***working group* links communities** from

observations, radiative transfer, transport, process & climate modelling

Observations / radiative transfer :

G. Stephens, H. Takahashi (NASA JPL), C. Stubenrauch, S. Protopapadaki, G. Sèze (LMD),
J. Luo, W. B. Rossow (CUNY), H. Masunaga (Nagoya Univ.), Roca (LEGOS), D. Bouniol (CNRM),
T. L'Ecuyer (Uni Wisconsin), S. Kato (NASA Langley), C. Schumacher (Texas Univ),
G. Mace, E. Zipser (Utah Univ), E. Jensen (NASA Ames), M. Krämer (FZ Jülich), A. Baran (MetOffice),
C. Kummerow (CSU), B. J. Sohn (Seoul Univ), H. Okamoto (Kyushu Univ)

Lagrangian transport, UTLS cirrus:

B. Legras, A.-S. Tissier, A. Hertzog (LMD)

Small scale process modelling :

S. van den Heever (CSU), R. Storer (NASA JPL), R. Plougonven, C. Muller (LMD),
W.-T. Chen (Nat Taiwan Univ), B. Kärcher (DLR)

Climate modelling :

T. Del Genio, G. Elsaesser (GISS), R. Ramaswamy, L. Donner (GFDL), B. Gasparini (ETHZ), U. Burkhardt (DLR), T. Mauritsen (MPI), M. Bonazzola, J.-B. Madeleine, C. Rio, C. Risi, S. Bony (LMD), R. Roehrig (CNRM)

GEWEX UTCC PROES highlights 2017

➤ 2nd GEWEX UTCC PROES meeting

hosted by Johnny Luo, at City College, University of New York, 28 – 29 March 2017

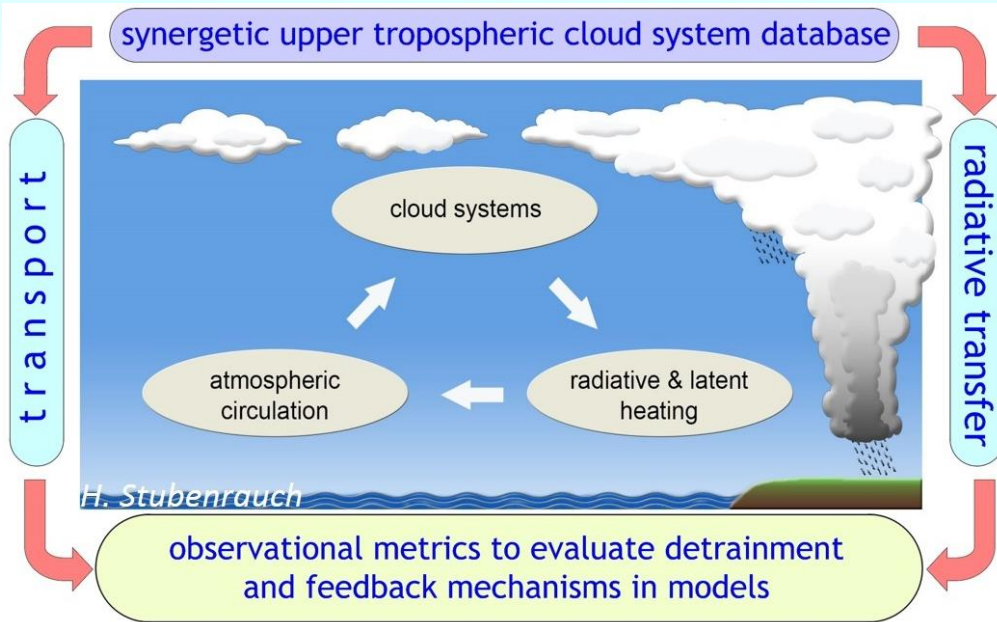
*vivid discussions about synergies & cooperations (funding dependent):
data, CRM studies & climate model parameterizations*



➤ article in GEWEX Newsletter May 2017

➤ **AGU session**, convened by H. Takahashi, R. Storer

➤ **build UTCC PROES website**, in cooperation with French data centre AERIS
goals, talks of the meetings, references



UTCC PROES Strategy

meetings: Nov 2015, Apr 2016, Mar 2017

working group links communities from observations, radiative transfer, transport, process & climate modelling

focus on tropical convective systems & cirrus originating from large-scale forcing

➤ **cloud system approach, anchored on IR sounder data**

horizontal extent & convective cores/cirrus anvil/thin cirrus ***based on*** p_{cld} , ϵ_{cld}

➤ **explore relationships between 'proxies' of convective strength & anvils**

➤ **build synergetic data** (vert. dimension, atmosph. environment, temporal res.)

➤ **determine heating rates** of different parts of UT cloud systems

➤ **follow snapshots** by Lagrangian transfer -> **evolution & feedbacks**

➤ **investigate how cloud systems behave in CRM studies**

& in GCM simulations (*under different parameterizations of convection/detrainment/microphysics*)

IR Sounders to derive UT cloud properties

TOVS, ATOVS

>1979 / ≥ 1995 : 7:30 / 1:30 AM/PM

AIRS, CrIS

≥ 2002 / ≥ 2012 : 1:30 AM/PM

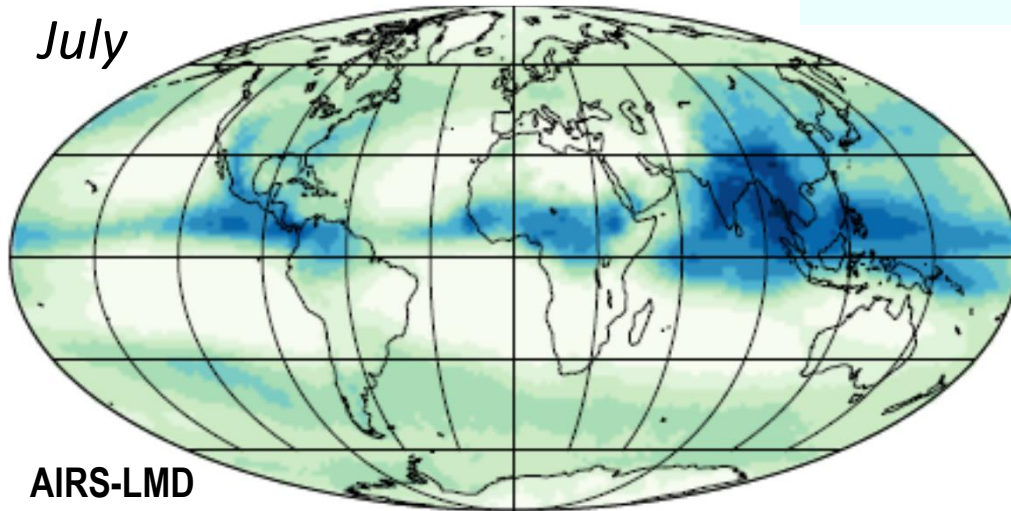
IASI (1,2,3), IASI-NG

≥ 2006 / ≥ 2012 / ≥ 2020 : 9:30 AM/PM

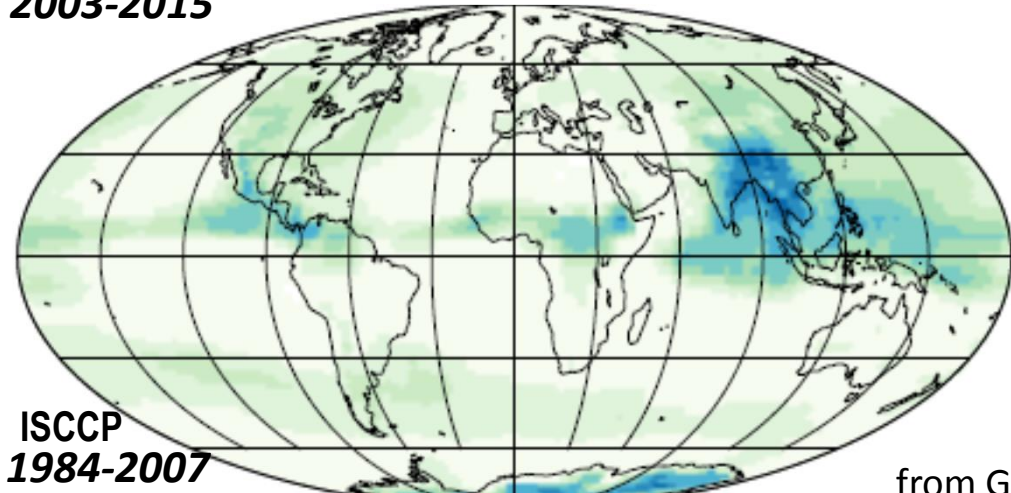
- long time series & good areal coverage
- good IR spectral resolution -> sensitive to cirrus
day & night, COD > 0.2, also above low clouds

UT cloud amount

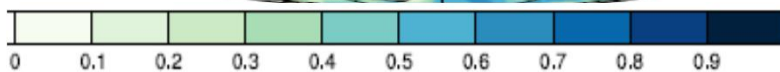
July



AIRS-LMD
2003-2015



ISCCP
1984-2007



CIRS (Cloud retrieval from IR Sounders):

Stubenrauch et al., J. Clim. 1999, 2006; ACP 2010, ACP 2017

AIRS / IASI climatologies -> French data centre AERIS

HIRS climatology -> EUMETSAT CM-SAF (DWD)

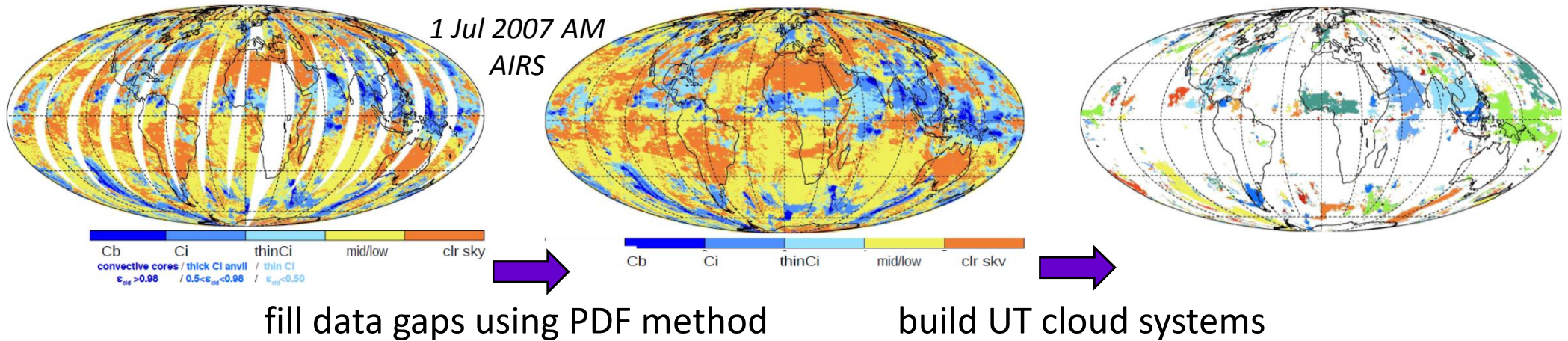
from GEWEX Cloud Assessment Database

<http://climserv.ipsl.polytechnique.fr/gewexca>

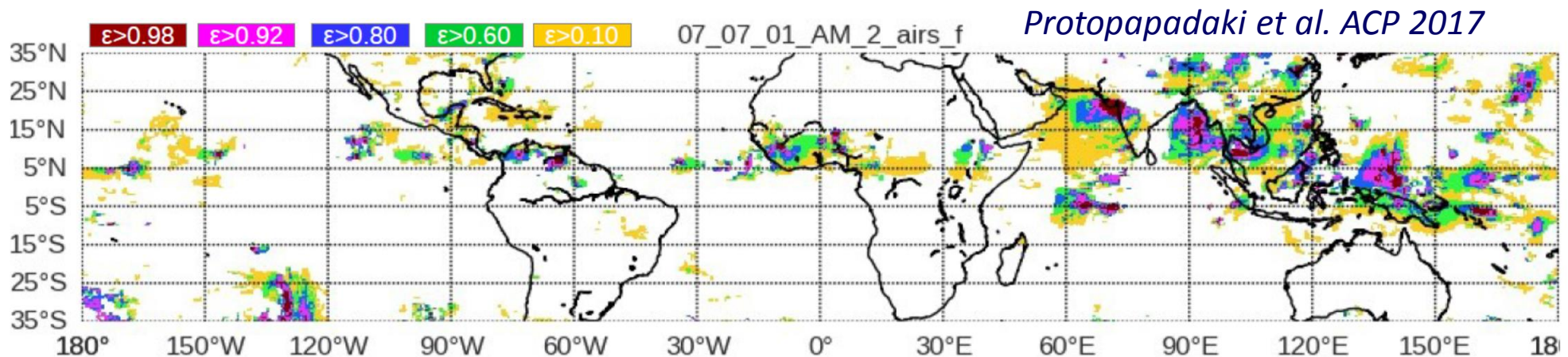
From cloud retrieval to cloud systems

clouds are extended objects, driven by dynamics -> organized systems

Method: 1) group adjacent grid boxes with high clouds of similar height (p_{cld})



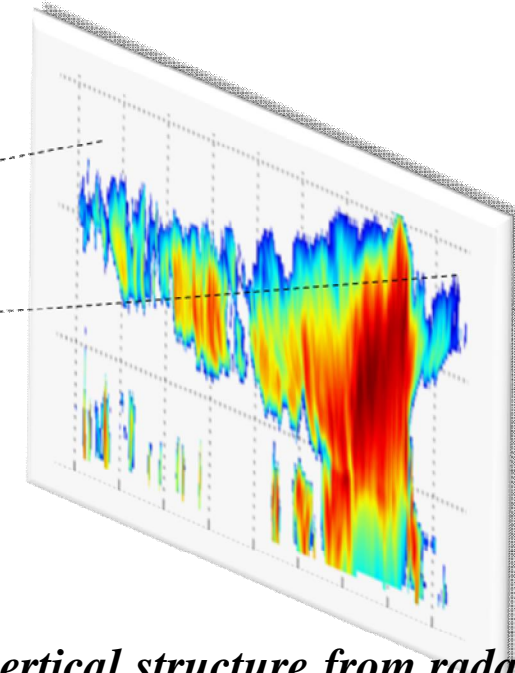
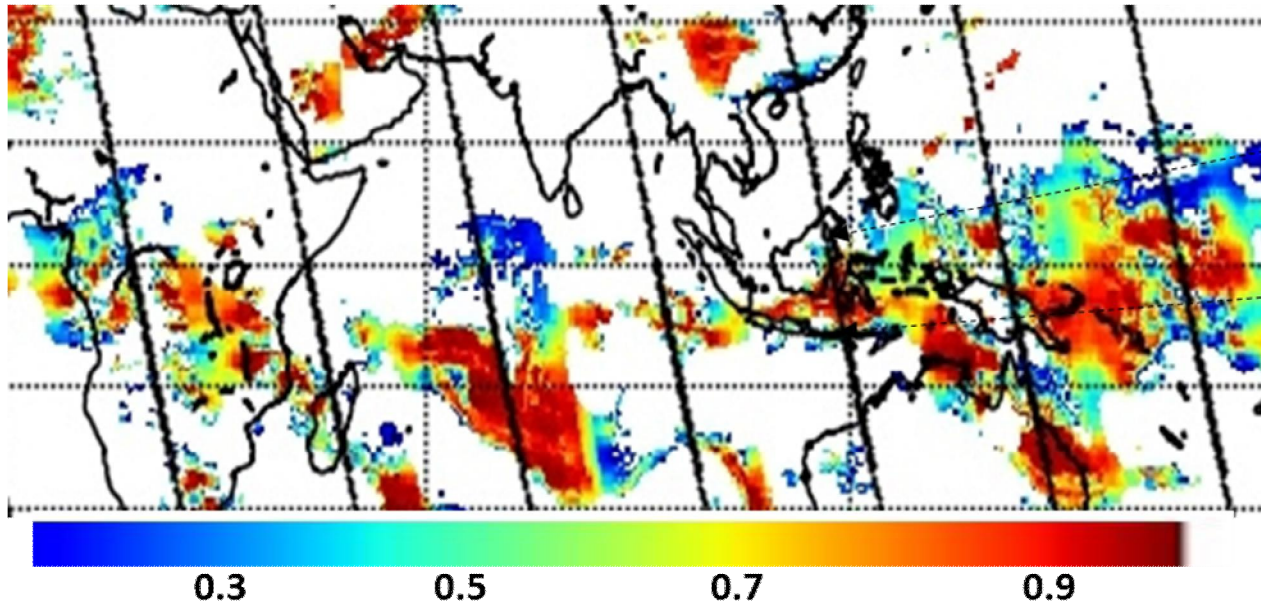
2) use ϵ_{cld} to distinguish convective core, thick cirrus, thin cirrus



30N-30S: UT cloud systems cover 20%, those without convective core 5%

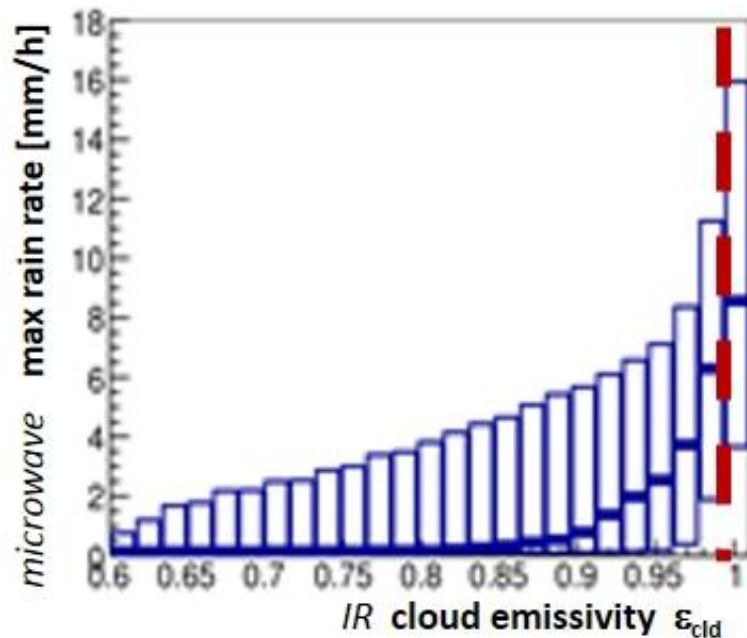
50% of these originate from convection (Luo & Rossow 2004, Riihimaki et al. 2012)

Observation synergies



horizontal emissivity structure of UT cloud systems
compared to other studies, IR sounders add thin cirrus ($0.1 < \epsilon_{cld} < 0.5$)

vertical structure from radar-lidar:
radar reflectivity of convective system



A-Train synergy (1h30 AM / PM)
AIRS – CALIPSO – CloudSat – AMSR-E

microwave imager - IR sounder synergy:
definition of convective core : $\epsilon_{cld} > 0.98$

Goal: relate anvil properties to convective strength

Strategy: need proxies

➤ to identify convective cores

$$\varepsilon_{cld} > 0.98 \quad (\text{compared to AMSR-E rain rate})$$

➤ to identify mature convective systems

$$\text{system core fraction} : 0.1 - 0.3 \quad (\text{reaching max core size})$$

➤ to describe convective strength

$$\text{core temp.} : T_{min}^{Cb} \quad (\text{Protopapadaki et al. 2017})$$

$$T_B^{IR} \quad (\text{Machado \& Rossow 1993})$$

vertical updraft : CloudSat Echo Top Height / TRMM

/ conv mass transport (Takahashi & Luo 2014 / Liu & Zipser 2007, Mullendore et al. 2008)

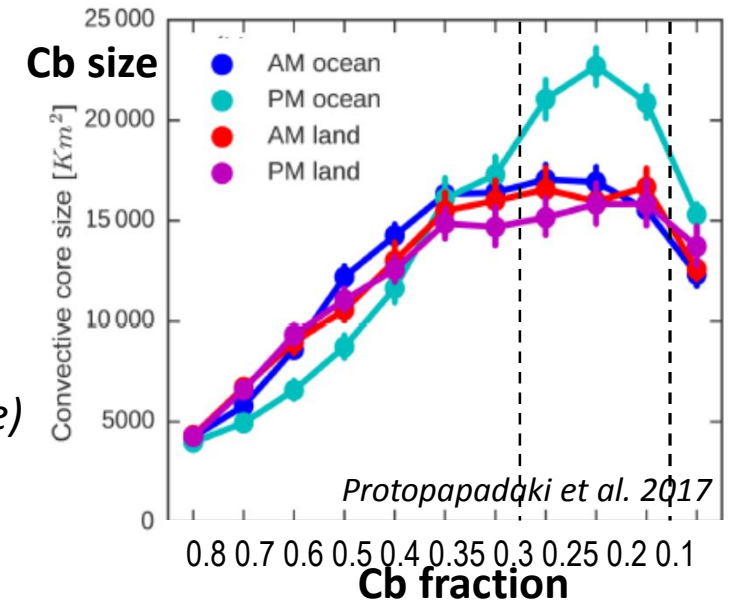
LNB : soundings / max mass flux outflow (Takahashi & Luo 2012)

heavy rain area: CloudSat-AMSR-E-MODIS (Yuan & Houze 2010)

core width : CloudSat (Igel et al. 2014)

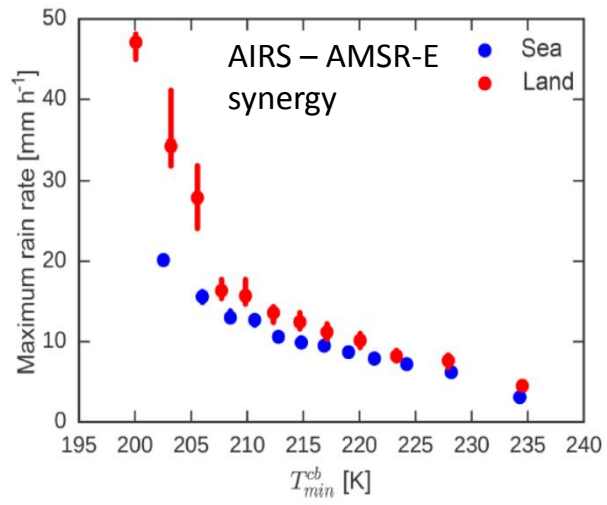
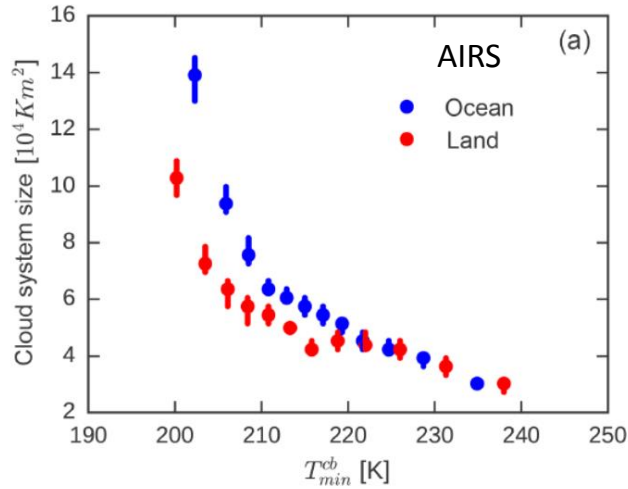
mass flux : ERA-Interim + Lagrangian approach (Tissier et al. 2016)

A-Train + 1D cld model (Masunaga & Luo 2016)



convective strength <-> cloud system properties

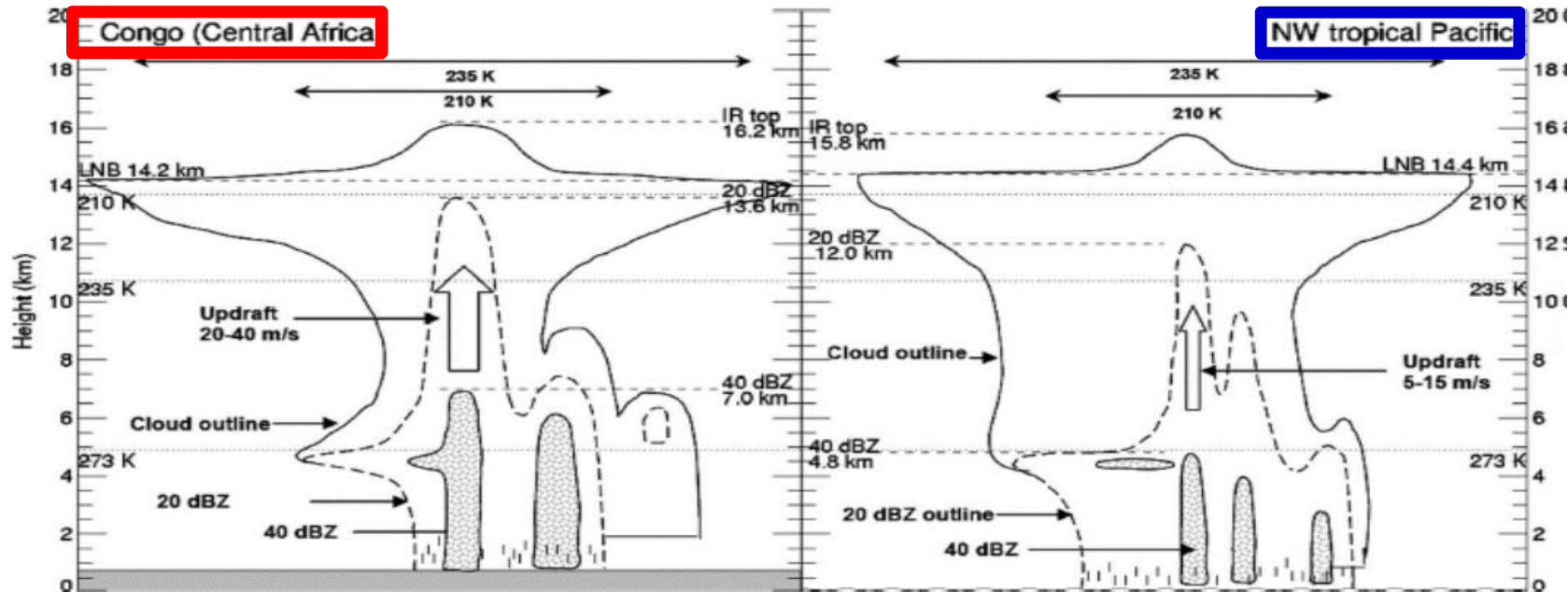
cloud system size / max rain rate increase with convective strength, but **land** – **ocean** differences



mature single convective core systems
Protopapadaki et al. 2017

Liu et al. 2007

typical strong convective systems (6-yr TRMM statistics)

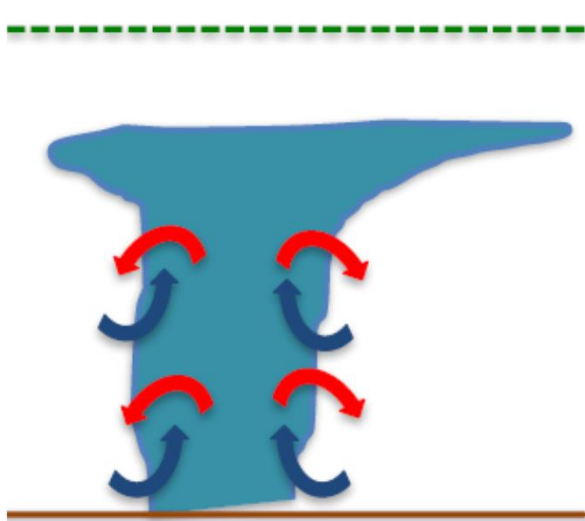


larger updraft & CC, smaller systems
less entrainment

smaller updraft & CC, larger systems
stronger entrainment *Takahashi et al. 2017*

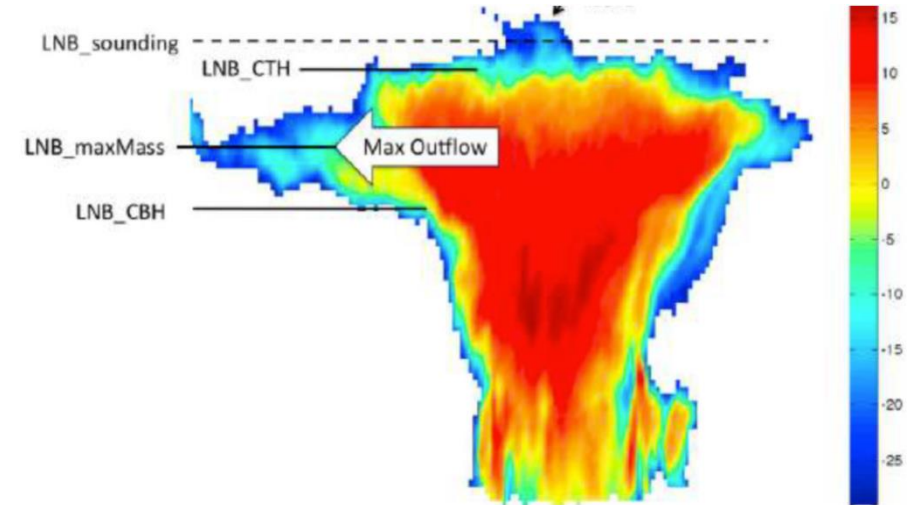
land -ocean convective entrainment

Takahashi et al. 2017



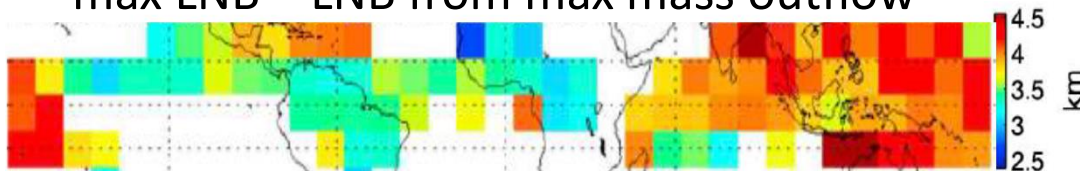
LNB from sounding:
max height
(no air mass interaction)

convective entrainment
affects buoyancy



convective cloud objects from 5-yr CloudSat statistics
Takahashi & Luo 2012

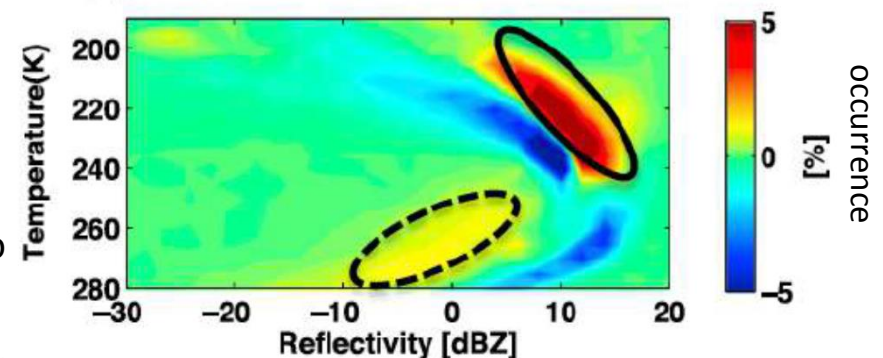
max LNB – LNB from max mass outflow



measure of magnitude of entrainment effect

more attenuation due to
heavy rain over land

(C) Africa & Amazonia – Warm Pool (CC)



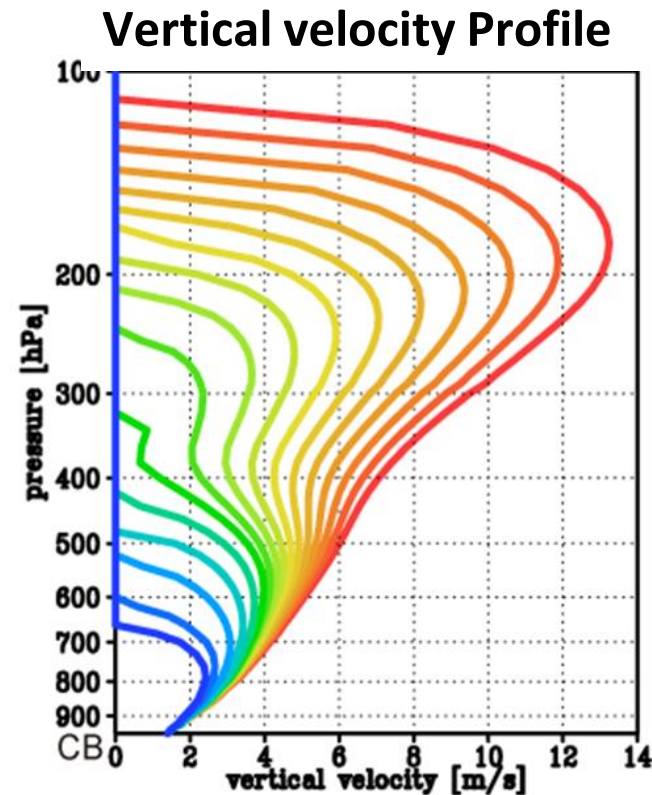
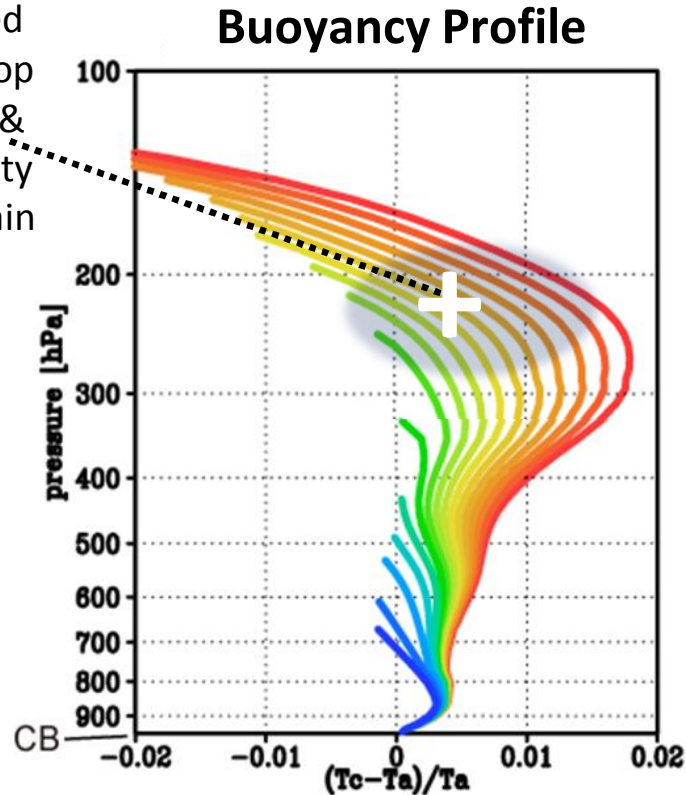
- Warm Pool more diluted than tropical land regions
- higher LNB(max mass outflow) associated with moister midtroposphere (reduced entrainment dilution) & smaller systems
- Tropical continental systems less entrainment => more intense convective cores

convective entrainment & vertical velocity

Masunaga & Luo 2016

Entraining ambient air will slow down convective updraft

Simulations constrained by cloud top buoyancy & vert velocity from A-Train synergy



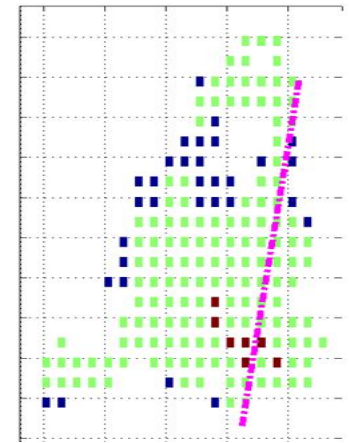
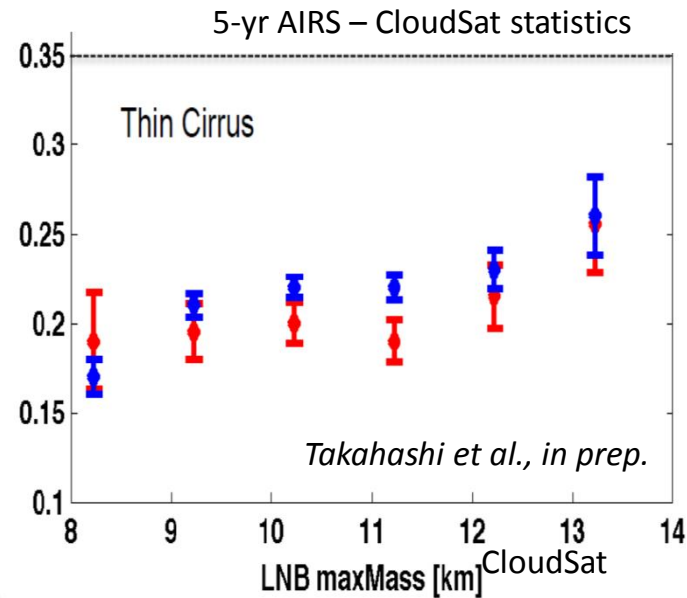
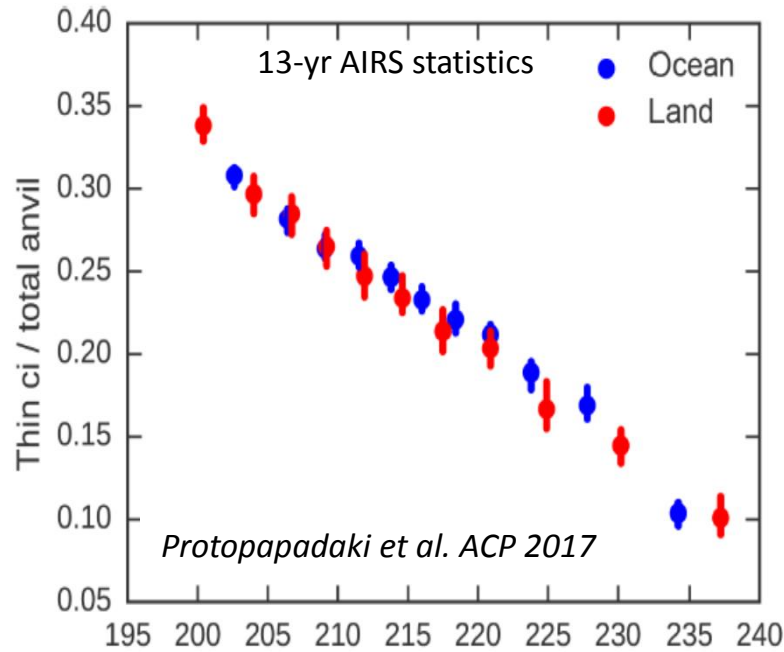
single column plume model

Plume entrainment rates from 0 km^{-1} to 0.4 km^{-1}

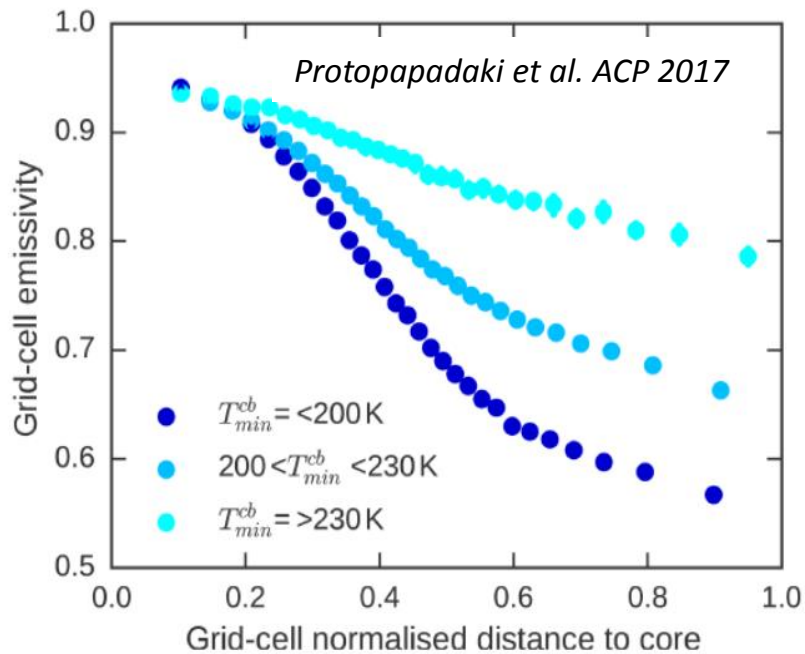
increasing entrainment leads to smaller climate sensitivity
(Zhao et al. 2014, Sanderson et al. 2010)
mechanisms not yet understood

L. Donner
UTCC PROES
meeting 2017

convective strength -> anvil properties



← T_{min}^{cb} [K] increasing convective strength →



Mature convective systems:

increase of thin Ci with increasing convective strength !
similar land / ocean
slope of ε_{cld} decrease with distance to core increases

relation robust using different proxies :

$$T_{min}^{Cb} / LNB(max\ mass)$$

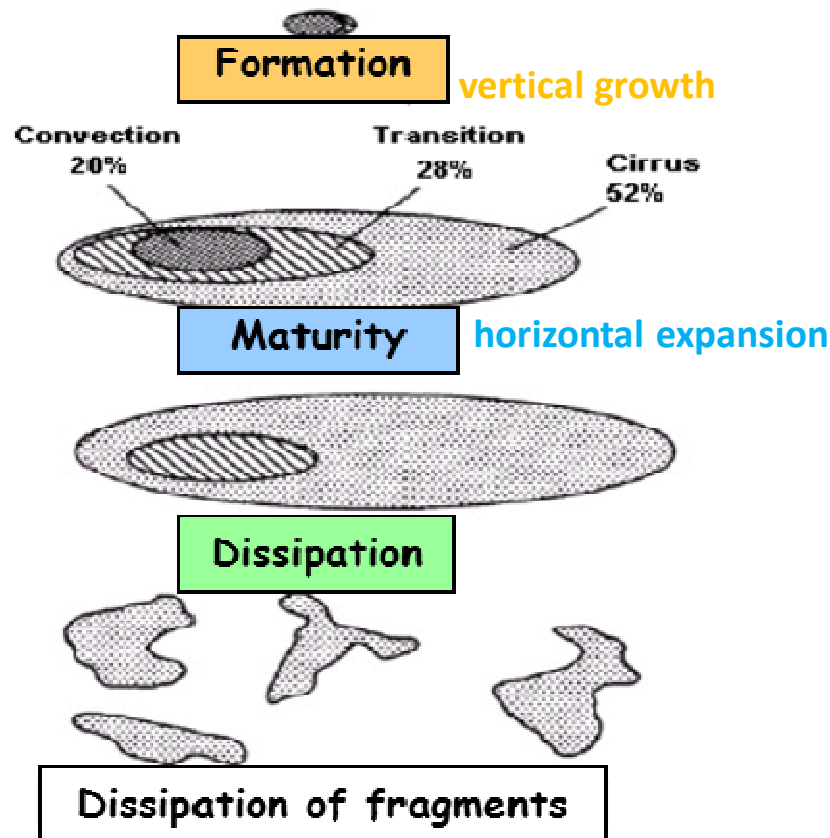
Life cycle of deep convective cloud systems

max of convection over **land** / **ocean** : **16-18h** / **early morning**

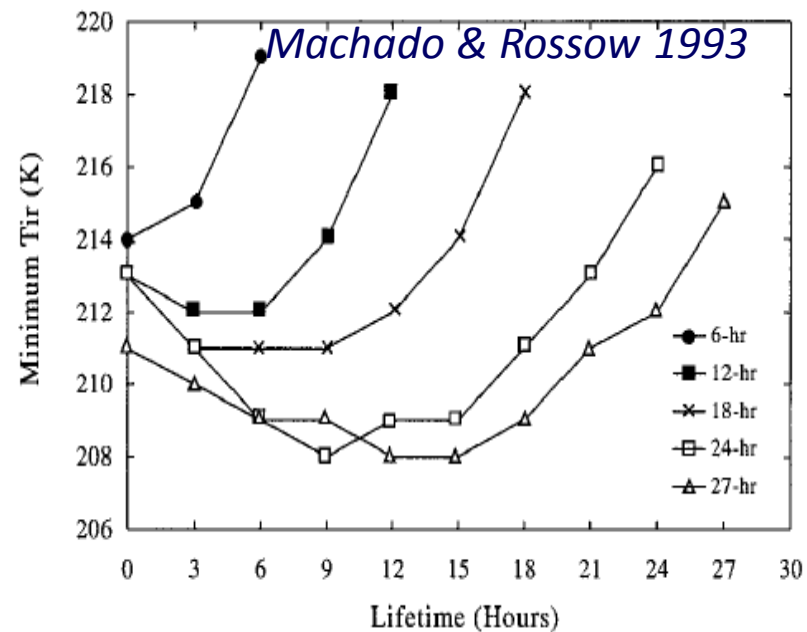
problem: most polar sunsynchronous observations do not catch this

-> use good time resolution of geostationary satellite imagers

& track cold convective cores with T_B^{IR} ; however T_B^{IR} depends on T_{cld} & on ϵ_{cld}



track all cold clouds ($T_B^{IR} < 245K$), sufficiently large (> 45 km) with ≥ 1 convective cloud (< 218 K)



coldest systems reach longest life-times

-> synergy AIRS/IASI + geostationary satellite imagers (for ex. HIMAWARI, G. Sèze)

Synergy with TRMM to analyze system life evolution

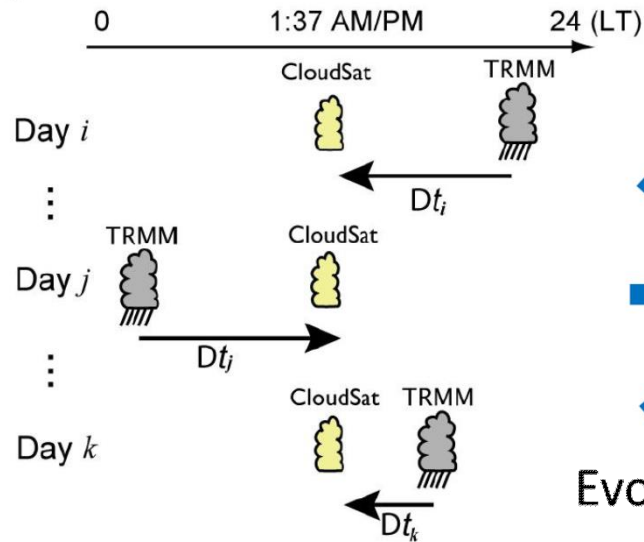
Composite observations w. r. t. convective life stages

H. Masunaga

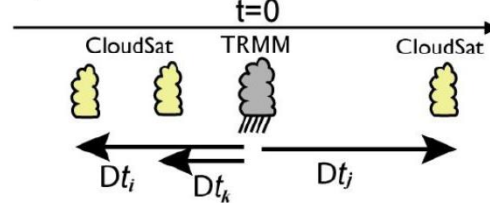
UTCC PROES meeting 2017

a) Instantaneous observations

20°S-20°N, ocean
2006-2009



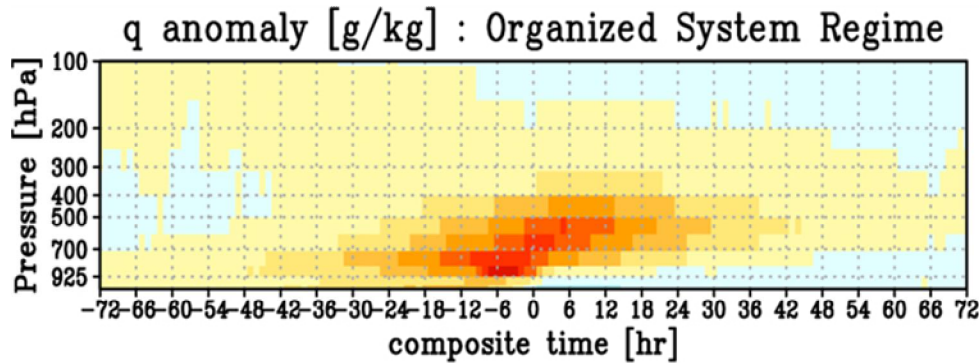
b) Composite time



Masunaga, 2012, 2013

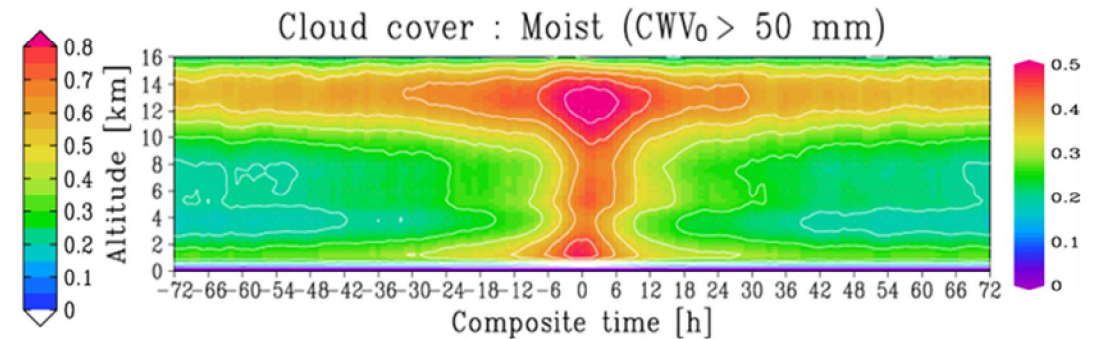
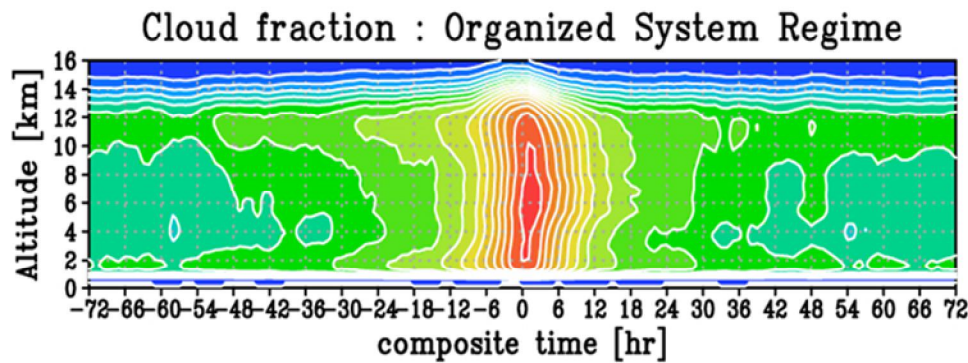
Masunaga & L'Ecuyer 2014

Evolution of moisture & cloud structures in organized convection



well defined convective cloud column at time of precipitation & then thinning out, but cirrus also around before convection
Radiative impact on convection ?

seminar at CSU



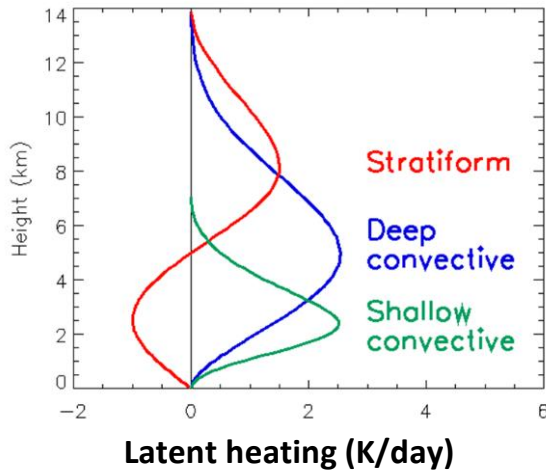
convective – anvil heating

latent (LH)

– radiative (RH)

*C. Schumacher
UTCC PROES
meeting 2017*

Schumacher et al. 2004

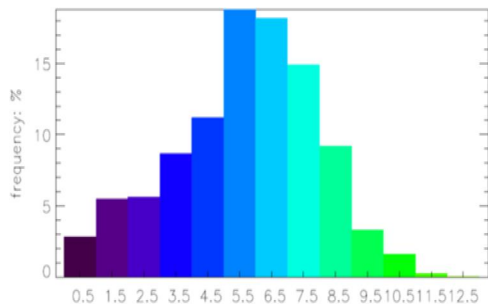


latent heating from TRMM :
column precipitation & cloud profile

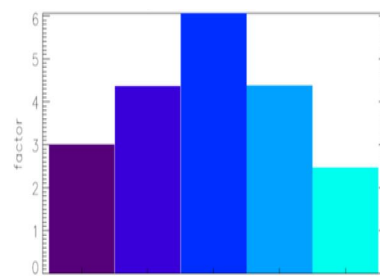
tropical stratiform rain leads to high peak in heating & cooling below
deep convective rain leads to broad atmospheric warming

Sensitivities of TRMM & CloudSat radar

Li & Schumacher 2010



depth of missed echo top (km)



echo base height (km)

TRMM radar misses 5 km to cloud top
& factor of 5 in horizontal extent

TRMM LH – ISCCP RH synergy

Li et al. 2013

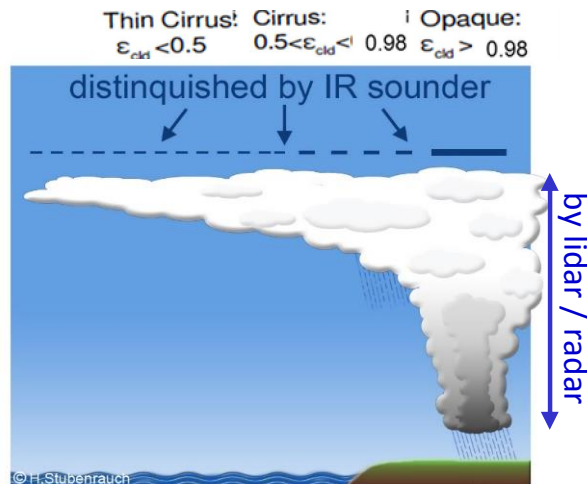
total radiative heating enhances gradient of latent heating at upper levels (e.g., 250 mb),
esp. over Africa, Maritime Continent & South America
& enhances overall LH by ~20%

heating rates of UT cloud systems (1)

critical to feedbacks : cirrus radiative heating in upper troposphere

➤ Cirrus anvils might regulate convection as they stabilize the atmospheric column by their heating (*Stephens et al. 2008, Lebsock et al. 2010*)

tropical convective regions: > 50% of total heating UT heating due to cirrus (*Sohn 1999*)
-> widespread impact on large-scale tropical atmospheric circulation



Heating will be affected by:

- areal coverage
- emissivity distribution
- vertical structure of cirrus anvils (layering & microphysics)

use nadir track info on vertical structure to propagate properties across UT cloud systems

- 1) assess existing radiative heating rates
- 2) sort FLXHR-LIDAR heating rates ...
- 3) compute heating rates ...

by categorizing cloud types wrt ϵ_{cld} & vertical structure
by categorizing atmospheric situation wrt T & H₂O profiles

Challenges:

- IWP, vertical profile of IWC
- ice crystal habit, size distribution -> SSP
- retrieval uncertainties in IWC / De profiles
- multiple cloud layering

Glance on actual heating rates

B. Legras, 2017

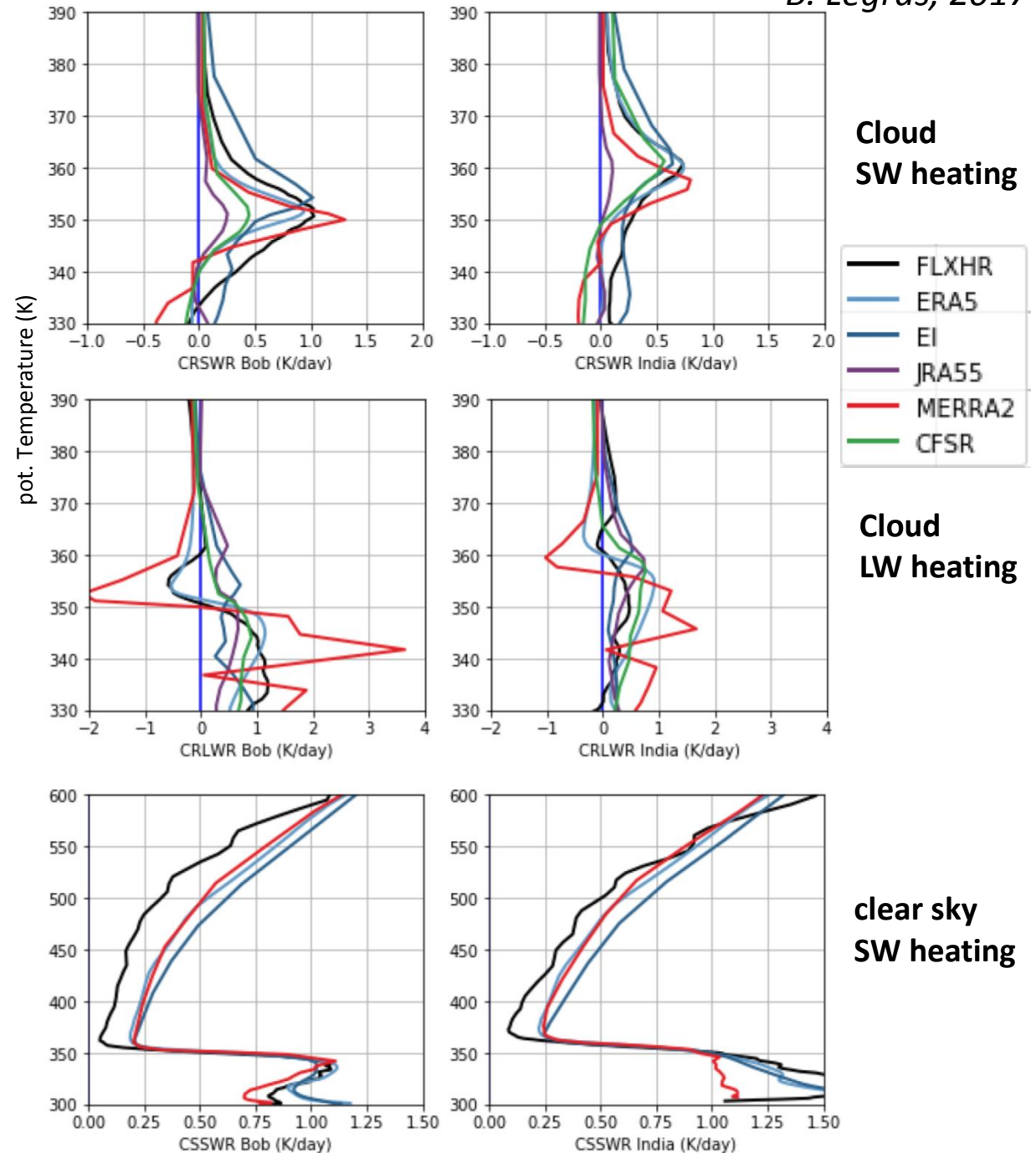


Reanalysis cloud heating rates disagree over convective regions, esp. over Asian monsoon region

Tissier & Legras ACP 2016

Obs. cloud heating rates disagree:
Johansson et al. 2015 (FLX-LIDAR)
 -> warming above clr sky LZRH
Yang et al. 2016 (CALIPSO)
 -> cooling above 16km

B. Legras, UTCC PROES meeting
 29 Apr 2016, Paris

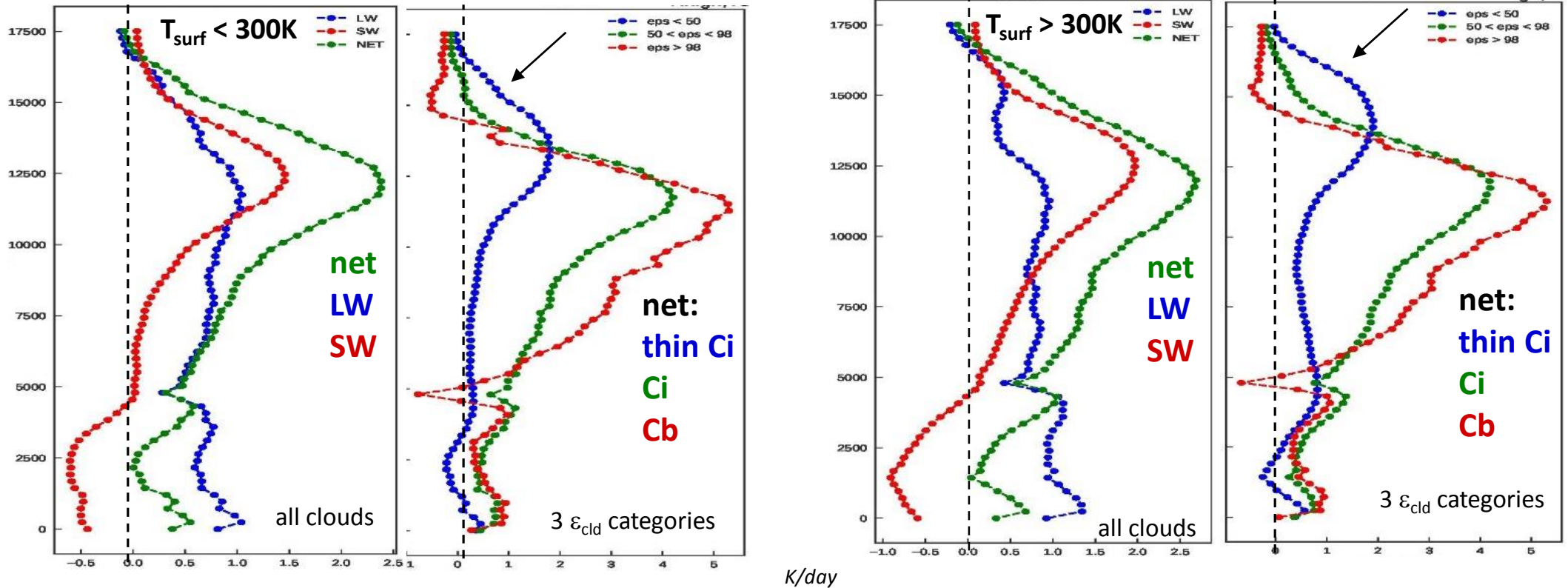


heating rates of UT cloud systems (3)

2) categorize Lidar-CloudSat FLXHR heating rates wrt to ϵ_{cld} , p_{cld} , vert. layering, thermodyn.

tropics, AIRS $p_{\text{cld}} < 200$ hPa, nadir track statistics

preliminary

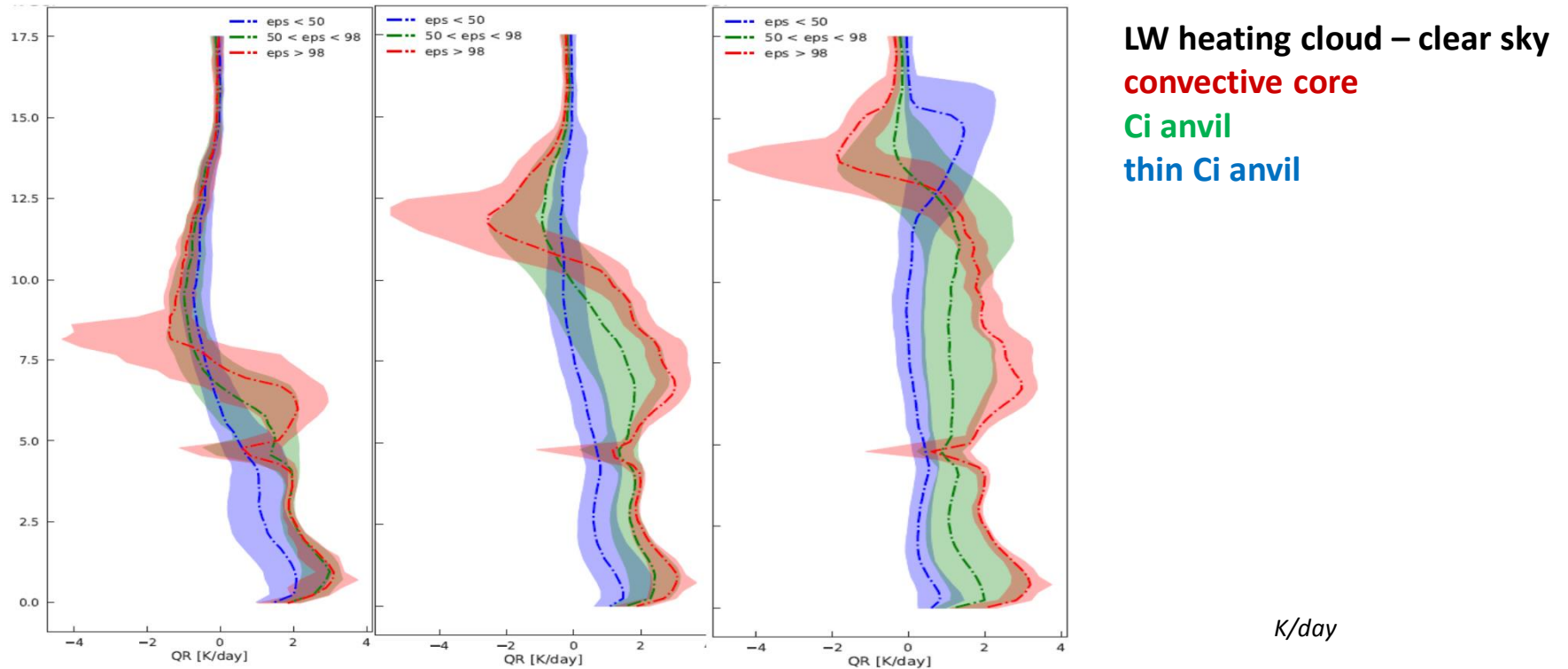


warmer T_{surf} \rightarrow UT cloud net heating occurring in thicker layers

heating rates of UT cloud systems (4)

AIRS UT cloud systems collocated to Lidar-CloudSat FLXHR heating rates wrt to ε_{cld} , p_{cld} ,

AIRS p_{cld} : 330 - 440 hPa p_{cld} : 200 - 330 hPa p_{cld} : 86 - 200 hPa *nadir track statistics* *preliminary*



clear distinction of heating associated with each category

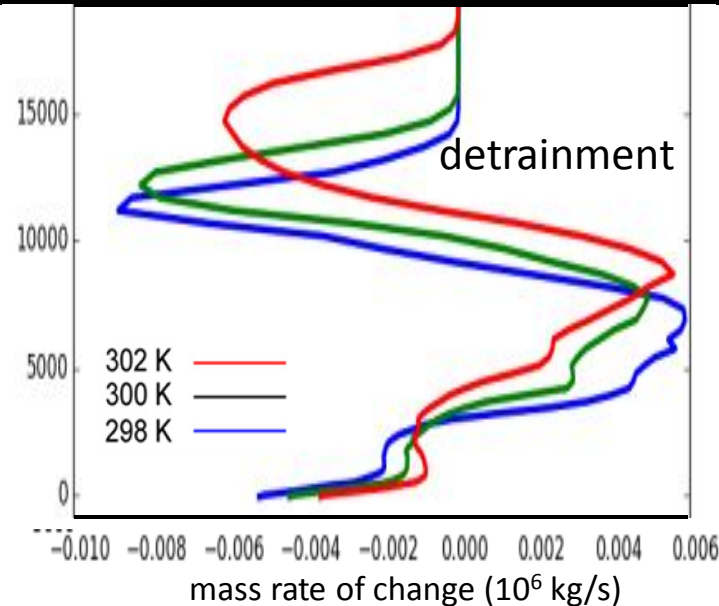
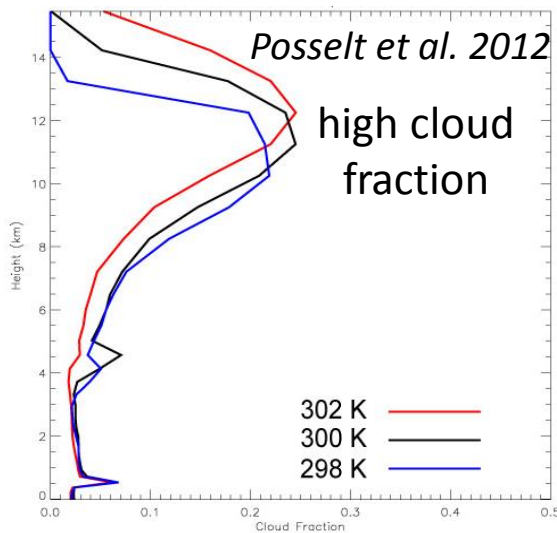
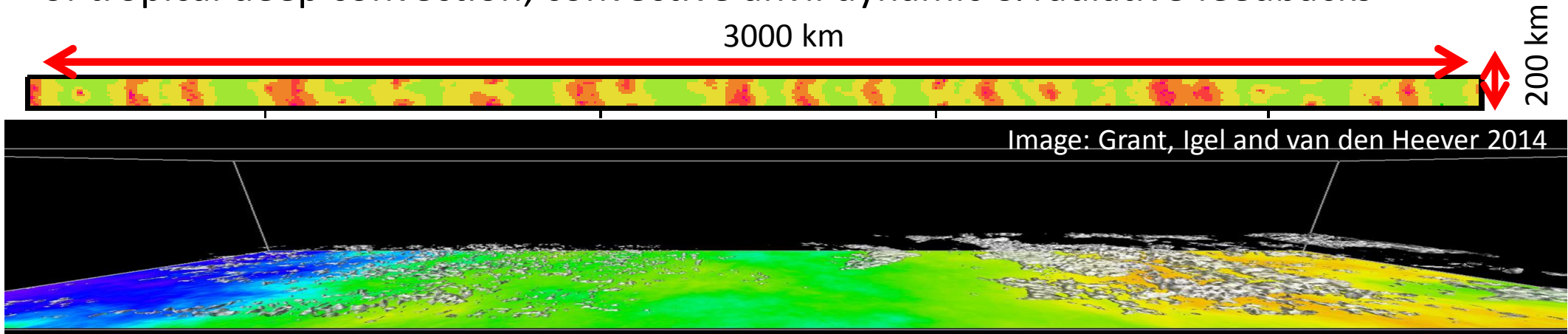
Next steps:

- refine categories, wrt atmospheric environment (meteorological reanalysis ERA5)
- expand heating rates across UT cloud systems

Characteristics of deep convection from CRM simulations

S. van den Heever, UTCC PROES meeting March 2017

advance our understanding of environmental impacts on horizontal & vertical scales of tropical deep convection; convective anvil dynamic & radiative feedbacks



Radiative-Convective Equilibrium simulations

R. Storer, water budget studies
UTCC PROES meeting
March 2017

detrainment higher & broader

increasing SST -> increased PW, convective intensity (w) & high cloud fraction, decrease in IR cooling -> slowing radiatively driven circulation

Challenge to simulate (organized) convection

A. Del Genio, G. Elsaesser

GEWEX PROES meeting March 2017

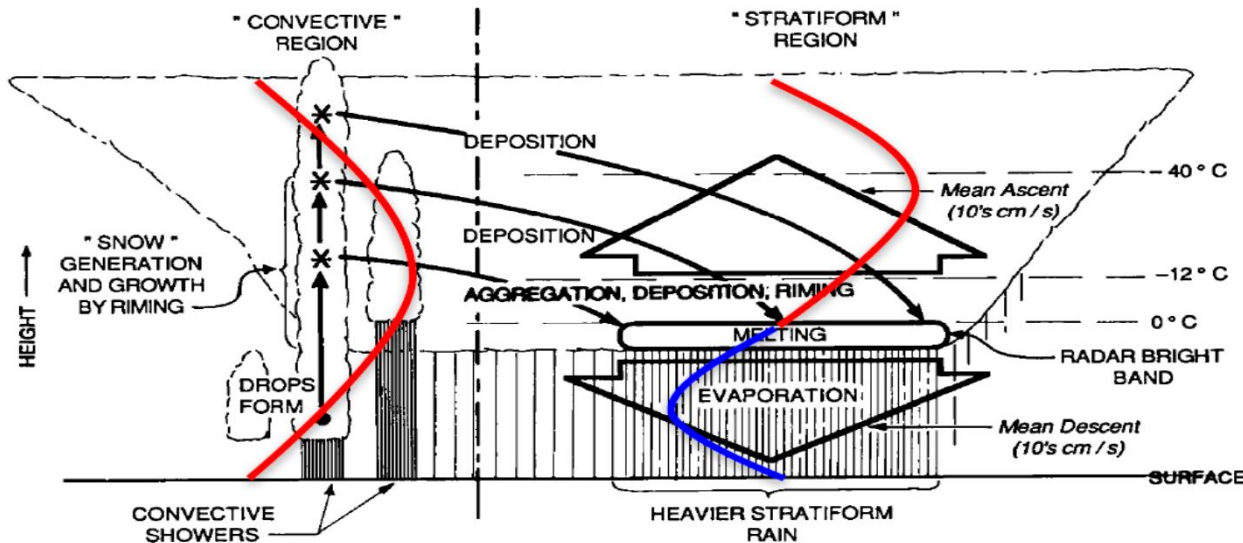


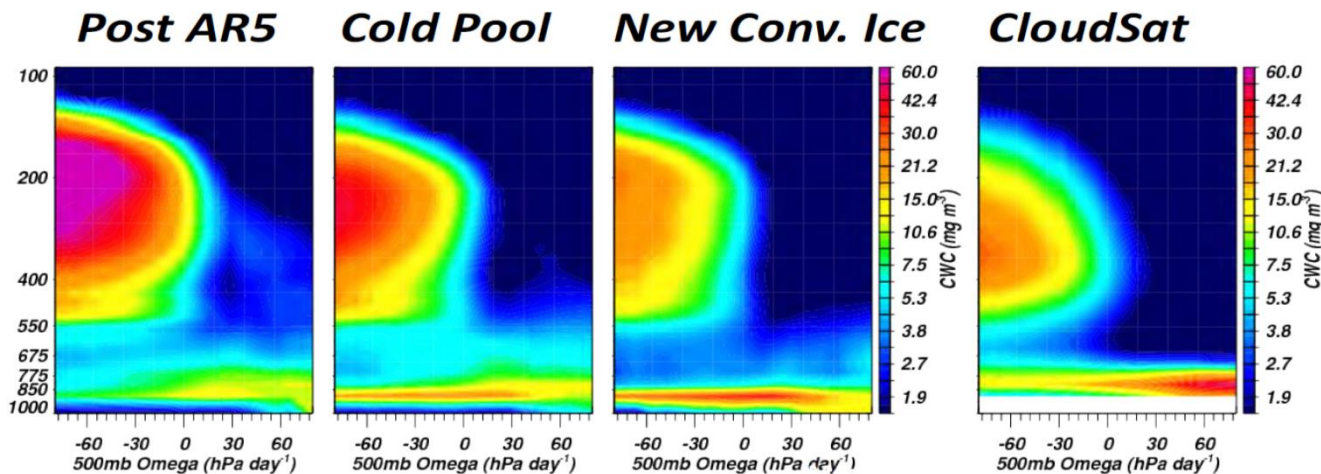
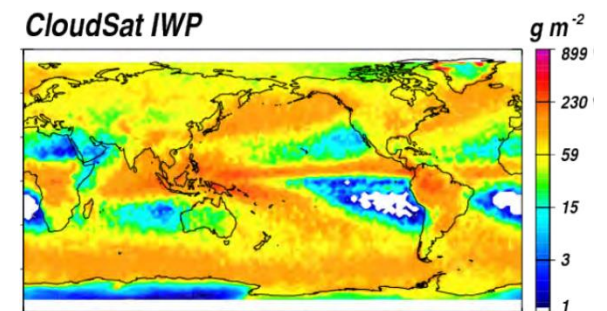
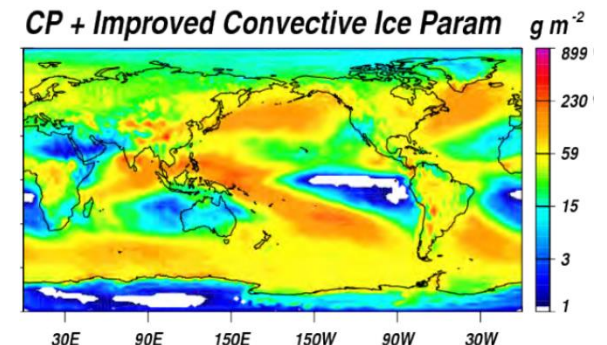
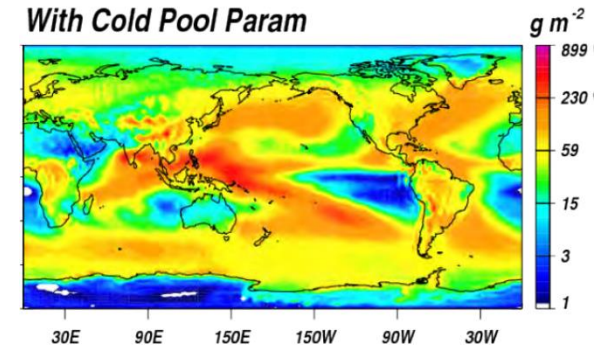
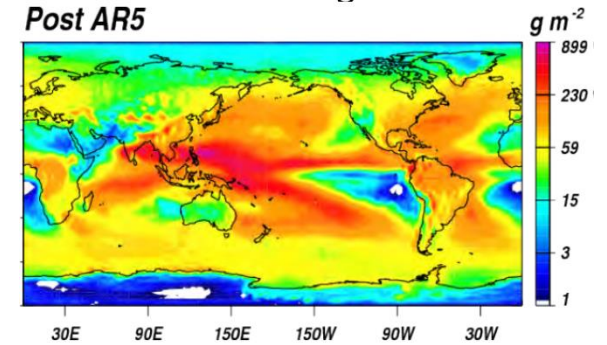
Figure 2. Schematic diagram of the precipitation mechanisms in a tropical cloud system. Solid arrows indicate particle trajectories (adapted from Houze 1989).

GISS GCM

Cold pools (Del Genio et al. 2015)

Convective ice parameterization (Elsaesser et al. 2017)

Microphysics (Morrison, Gettelman 2015)



need to simulate entire convective life cycle...

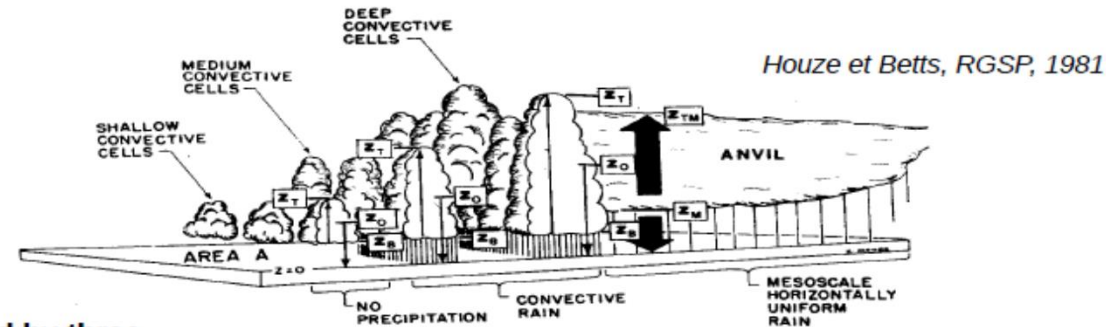
Convection & UT clouds in LMDZ climate model

convection

C. Rio

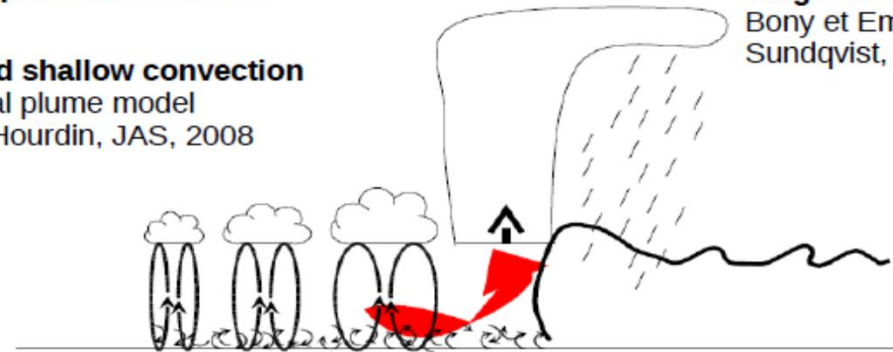
UTCC PROES

meeting 2015



Handled by three different parameterizations

Dry and shallow convection
Thermal plume model
Rio et Hourdin, JAS, 2008



Large-scale condensation
Bony et Emanuel, JAS, 2001
Sundqvist, QJRMS, 1978

Deep convection and associated cold pools
Emanuel, JAS, 1991 revisited by
Grandpeix et Lafore, JAS, 2010

- (1) Formed by large-scale advection and deep convection (anvils) ; in this latter case they

depend on the detrainment of water vapor and maximum precipitation efficiency ϵ_{\max}

[Emanuel & Živković-Rothman 1999 ; Bony & Emanuel 2001; Rio et al. 2012 ; Grandpeix & Lafore 2010]

- (2) Phase based on temperature using $x_{\text{liq}} = \left(\frac{T - T_{\text{ice}}}{T_0 - T_{\text{ice}}} \right)^{n_x}$

- (3) Precipitation mass flux ($\rho w_{iw} q_{iw}$) computed using ice particle fall velocity

$$w_{iw} = \gamma_{iw} w_0 \quad \text{with } w_0 = 3.29(\rho q_{iw})^{0.16} \quad \text{and } \gamma_{iw} \text{ a tuning coefficient}$$

[Zender and Kiehl, 1997 ; Heymsfield and Donner, 1990]

UT clouds

J.-B. Madeleine

UTCC PROES

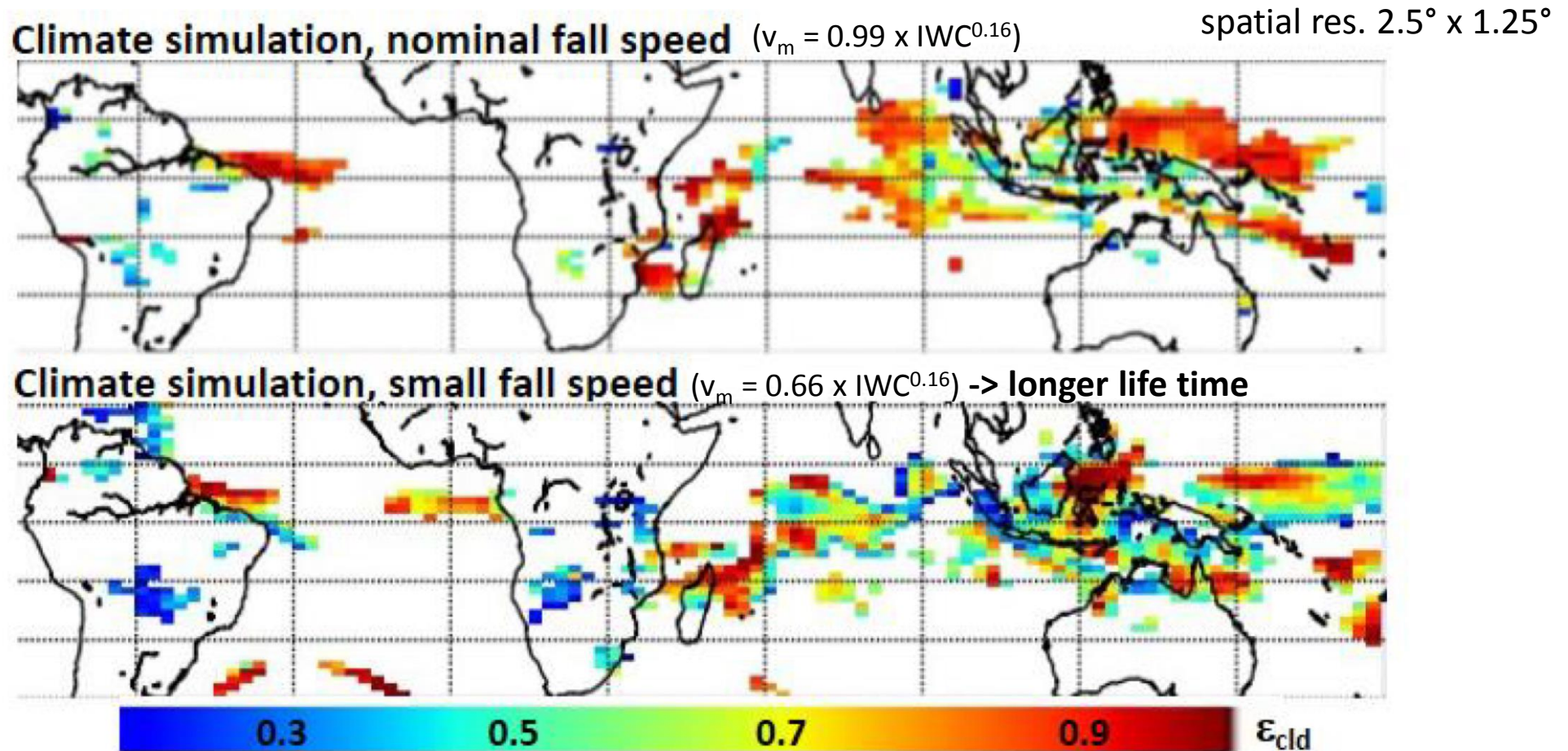
meeting 2015

Diagnostics for UT cloud assessment in climate models

M. Bonazzola, C. Stubenrauch, S. Protopapadaki

*analyze GCM clouds as seen from AIRS/IASI, via simulator
& construct UT cloud systems*

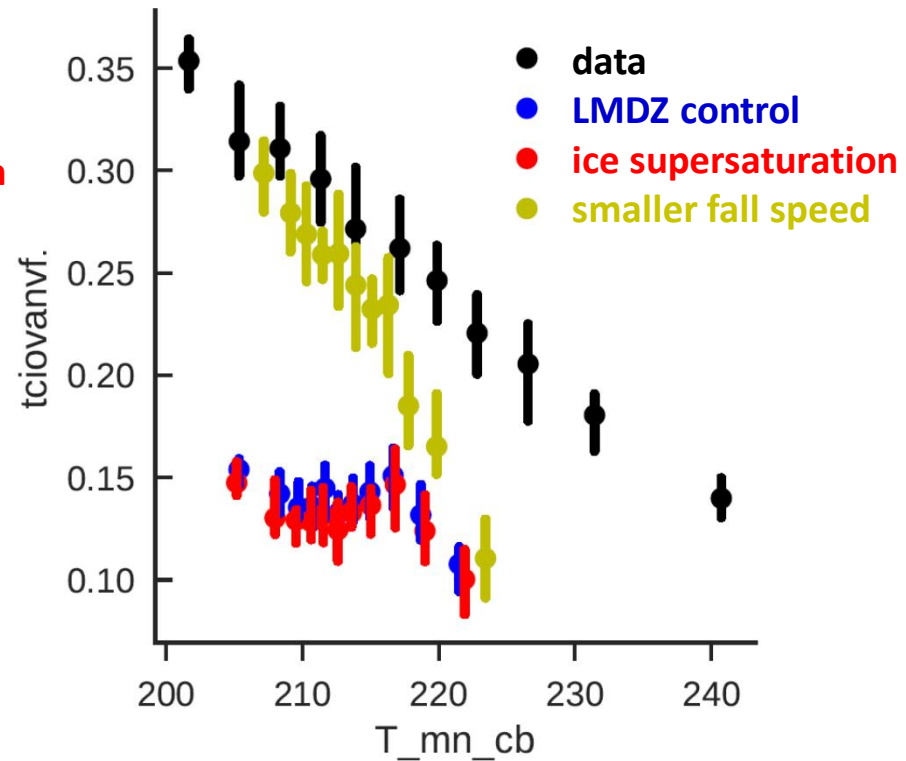
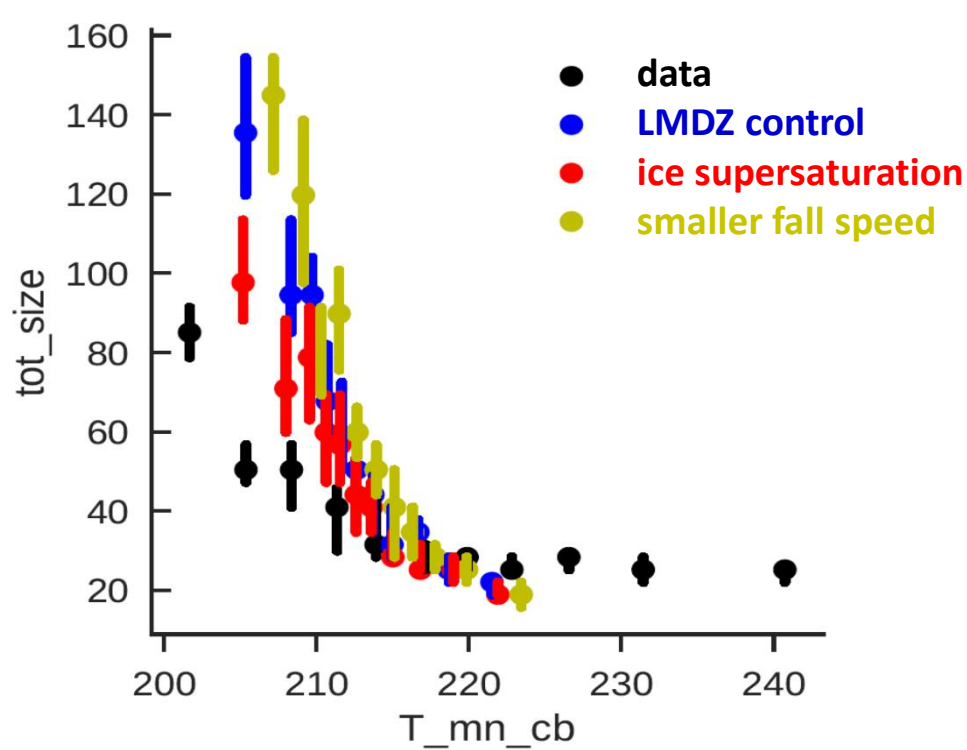
-> evaluation of GCM convection schemes / detrainment / microphysics
allows to assess horizontal extent & emissivity structure of UT cloud systems



horizontal cloud system emissivity structure sensitive to fall speed

observational metrics to probe process understanding

How do the anvil properties change with convective strength?



➤ **UT cloud system size increases with convective strength**

*LMDZ cloud system size increases stronger than obs,
including ISS leads to smaller systems, in better agreement to obs*

➤ **thin Ci over anvil area increases with convective strength**

*thin Ci over anvil area much too small for LMDZ and no correlation between 205-215K
reducing the fall speed (longer life time of cirrus) leads to an increase with convective strength*

Summary & Outlook

working group (*meetings: Nov 2015, Apr 2016, Mar 2017 -> vivid discussions*)

cooperations being formed (depending on funding), focus on tropical convective systems

- **synergetic cloud system data based on IR sounder data**
powerful tool to study relation between convection & anvil properties
- **relation between convective strength & anvil properties:**
change in emissivity structure -> can be used to constrain models
- **classification of vertical structure & heating rates (A-Train synergy)**
-> extend to UT cloud systems & integrate into feedback studies
using Lagrangian transport & advanced analysis methods
- **investigate how cloud systems behave in CRM studies**
& in GCM simulations (*under different parameterizations of convection/detrainment/microphysics*)
- **assessment of heating rates ?**

next meeting: autumn 2018 (Paris)